

Shallow water theories

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Shallow water

Long wave assumption:

- Typical wavelength L is large compared to the mean depth d , i.e., $d/L \ll 1$.
- Typical period T is large too because phase velocities are closed to the critical speed $c_0 = \sqrt{gd}$, i.e., $L/T \approx c_0$.

Distorsion:

$$\xi = \epsilon x, \quad \tau = \epsilon t \quad \Rightarrow \quad \frac{\partial}{\partial x} = \epsilon \frac{\partial}{\partial \xi}, \quad \frac{\partial}{\partial t} = \epsilon \frac{\partial}{\partial \tau}.$$

Shallow water equations

$$\epsilon^2 \frac{\partial^2 \phi}{\partial \xi^2} + \frac{\partial^2 \phi}{\partial y^2} = 0 \quad -d \leq y \leq \eta,$$

$$\frac{\partial \phi}{\partial y} = 0 \quad y = -d,$$

$$g\eta + \epsilon \frac{\partial \phi}{\partial \tau} + \frac{\epsilon^2}{2} \left(\frac{\partial \phi}{\partial \xi} \right)^2 + \frac{1}{2} \left(\frac{\partial \phi}{\partial y} \right)^2 = 0 \quad y = \eta,$$

$$\frac{\partial \phi}{\partial y} - \epsilon \frac{\partial \eta}{\partial \tau} - \epsilon^2 \frac{\partial \phi}{\partial \xi} \frac{\partial \eta}{\partial \xi} = 0 \quad y = \eta.$$

Small parameter expansion:

$$\phi = \epsilon \phi_1 + \epsilon^3 \phi_3 + \dots, \quad \eta = \epsilon^2 \eta_2 + \epsilon^4 \eta_4 + \dots.$$

Resolution at orders 1 and 2

Laplace equation and bottom impermeability at order ϵ :

$$\begin{aligned}\frac{\partial^2 \phi_1}{\partial y^2} &= 0, & \frac{\partial \phi_1}{\partial y} \Big|_{y=-d} &= 0, \\ \Rightarrow \phi_1 &= \Phi_1(\xi, \tau).\end{aligned}$$

Surface isobarity at order ϵ^2 :

$$g \eta_2 + \frac{\partial \phi_1}{\partial \tau} \Big|_{y=0} = 0 \quad \Rightarrow \quad \eta_2 = -\frac{1}{g} \frac{\partial \Phi_1}{\partial \tau}.$$

Resolution at order 3

Laplace equation and bottom impermeability at order ϵ^3 :

$$\frac{\partial^2 \phi_1}{\partial \xi^2} + \frac{\partial^2 \phi_3}{\partial y^2} = 0, \quad \left. \frac{\partial \phi_3}{\partial y} \right|_{y=-d} = 0,$$
$$\Rightarrow \phi_3 = \Phi_3(\xi, \tau) - \frac{(y+d)^2}{2} \frac{\partial^2 \Phi_1}{\partial \xi^2}.$$

Surface impermeability at order ϵ^3 :

$$\left. \frac{\partial \phi_3}{\partial y} \right|_{y=0} = \frac{\partial \eta_2}{\partial \tau} \quad \Rightarrow \quad \frac{\partial^2 \Phi_1}{\partial \tau^2} - g d \frac{\partial^2 \Phi_1}{\partial \xi^2} = 0,$$
$$\Rightarrow \Phi_1 = \Phi_1^-(\xi - c\tau) + \Phi_1^+(\xi + c\tau), \quad c \equiv \sqrt{gd}.$$

Special case:

$$\Phi_1 = \Phi_1(\xi - c\tau).$$

Resolution at order 4 and 5

Surface isobarity at order ϵ^4 :

$$g \eta_4 + \left[\frac{\partial \phi_3}{\partial \tau} + \frac{1}{2} \left(\frac{\partial \phi_1}{\partial \xi} \right)^2 \right]_{y=0} = 0$$
$$\Rightarrow \eta_4 = -\frac{1}{g} \frac{\partial \Phi_3}{\partial \tau} + \frac{d^2}{2g} \frac{\partial^3 \Phi_1}{\partial^2 \xi \partial \tau} - \frac{1}{2g} \left(\frac{\partial \Phi_1}{\partial \xi} \right)^2.$$

Laplace equation and bottom impermeability at order ϵ^5 :

$$\frac{\partial^2 \phi_3}{\partial \xi^2} + \frac{\partial^2 \phi_5}{\partial y^2} = 0, \quad \frac{\partial \phi_5}{\partial y} \Big|_{y=-d} = 0,$$
$$\Rightarrow \phi_5 = \Phi_5(\xi, \tau) - \frac{(y+d)^2}{2} \frac{\partial^2 \Phi_3}{\partial \xi^2} + \frac{(y+d)^4}{4!} \frac{\partial^4 \Phi_1}{\partial \xi^4}.$$

Surface impermeability at order ϵ^5 :

$$\left. \frac{\partial \phi_5}{\partial y} \right|_{y=0} + \eta_2 \left. \frac{\partial \phi_3}{\partial y} \right|_{y=0} = \frac{\partial \eta_4}{\partial \tau} + \frac{\partial \eta_2}{\partial \xi} \left. \frac{\partial \phi_3}{\partial \xi} \right|_{y=0}.$$

Special case $\Phi_n = \Phi_n(\theta)$ with $\theta \equiv \xi - c\tau$:

$$\frac{c}{6} \frac{d^2}{d\theta^4} \Phi_1 + \frac{3}{2} \frac{d\Phi_1}{d\theta} \frac{d^2 \Phi_1}{d\theta^2} = 0,$$

or with $\eta_2 = (c/g) d\Phi_1/d\theta$:

$$\frac{c}{6} \frac{d^2}{d\theta^3} \eta_2 + \frac{3g}{2c} \eta_2 \frac{d\eta_2}{d\theta} = 0.$$

General resolution for travelling waves

First integration:

$$\frac{c d^2}{6} \frac{d^2 \eta_2}{d\theta^2} + \frac{3g}{4c} \eta_2^2 = \text{constant} \equiv K_1.$$

Second integration (after multiplication by η_2'):

$$\frac{c d^2}{12} \left(\frac{d \eta_2}{d\theta} \right)^2 + \frac{g}{4c} \eta_2^3 = K_1 \eta_2 + K_2.$$

General solution via Jacobian elliptic functions.

Solitary wave

Ansatz: $\eta_2 = \alpha_0 + \alpha_1 \operatorname{sech}^2(\alpha_2 \theta)$.

After substitution into the equation:

$$\alpha_2 = \frac{\sqrt{3 g \alpha_1}}{2 d c}, \quad \alpha_0 = -\frac{\alpha_1}{3}, \quad K_1 = \frac{g \alpha_1^2}{12 c}, \quad K_2 = \frac{g \alpha_1^3}{54 c}.$$

Problem: $\alpha_0 = K_1 = K_2 = 0$ only if $\alpha_1 = 0$, that is for zero amplitude!?

Remedy: renormalisation. Observations:

- d is **not** the mean depth!
- c is **not** the phase velocity!

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Approximate solution

Free surface:

$$\eta \approx \epsilon^2 \eta_2 = \epsilon^2 \alpha_0 + \epsilon^2 \alpha_1 \operatorname{sech}^2(\alpha_2 \theta).$$

Horizontal velocity:

$$u \approx \epsilon^2 \frac{d\Phi_1}{d\theta} = \epsilon^2 \frac{g \eta_2}{c} \approx \frac{g \eta}{c}.$$

Physical parameters

Total local water depth: $h \equiv d + \eta$.

– Mean physical (or real) depth d^* for a solitary wave such that: $h(\infty) = d^*$. Thence:

$$d^* = d + \epsilon^2 \alpha_0 \quad \Rightarrow \quad d = d^* + \epsilon^2 \alpha_1 / 3.$$

– Free surface elevation from rest:

$$\eta^* \equiv h(\theta) - h(\infty) \approx \epsilon^2 \alpha_1 \operatorname{sech}^2(\alpha_2 \theta).$$

– Wave amplitude $a \equiv \eta^*(0)$, thence $a = \epsilon^2 \alpha_1$.

Apparent phase velocity:

$$c = \sqrt{gd} = \sqrt{g(d^* + \epsilon^2 \alpha_1/3)} \approx \sqrt{gd^*} \left(1 + \frac{\epsilon^2 \alpha_1}{6d^*} \right).$$

'Fixed' reference frame: $u^* = u - u(\infty)$.

Galilean transformation:

$$\begin{aligned} \theta^* &= \theta + u(\infty)\tau = \xi - c\tau + \epsilon^2 (g/c) \alpha_0 \tau \\ &= \xi - \left(c + \epsilon^2 \frac{g}{3c} \alpha_1 \right) \tau. \end{aligned}$$

Phase velocity in the 'fixed' frame of reference:

$$\begin{aligned} c^* &\equiv c + \epsilon^2 \frac{g}{3c} \alpha_1 \approx \sqrt{gd^*} \left(1 + \frac{\epsilon^2 \alpha_1}{2d^*} \right) \\ &= \sqrt{gd^*} \left(1 + \frac{a}{2d^*} \right). \end{aligned}$$

Solitary wave characteristics

Froude number:

$$\text{Fr} \equiv \frac{c^*}{\sqrt{g d^*}} \approx 1 + \frac{a}{2 d^*} \geq 1.$$

Solitary wave speed is super-critical.

Summary

Approximations break symmetries, in general.

Parameters may not be what they seem to be.

(Functions and operators may have the same 'problem'.)

A posteriori check of parameters definition.

(Parameters must be well defined!)

Renormalise, if needed.

(Removal of secular terms. We must know/understand the Physics!)

Other techniques: Multiple scales, Lie group transformations, etc.

Multiple scale variant

Distorsion for almost unidirectional waves:

$$\xi = \epsilon(x - c_0 t), \quad \tau = \epsilon^3 t, \quad c_0 = \sqrt{gd},$$
$$\Rightarrow \quad \frac{\partial}{\partial x} = \epsilon \frac{\partial}{\partial \xi}, \quad \frac{\partial}{\partial t} = \epsilon^3 \frac{\partial}{\partial \tau} - \epsilon c_0 \frac{\partial}{\partial \xi}.$$

Small parameter expansion:

$$\phi = \epsilon \phi_1 + \epsilon^3 \phi_3 + \dots, \quad \eta = \epsilon^2 \eta_2 + \epsilon^4 \eta_4 + \dots.$$

Solution of Laplace and bottom equations:

$$\phi_1 = \Phi_1(\xi, \tau), \quad \phi_3 = \Phi_3(\xi, \tau) - \frac{(y+d)^2}{2} \frac{\partial^2 \Phi_1}{\partial \xi^2},$$
$$\phi_5 = \Phi_5(\xi, \tau) - \frac{(y+d)^2}{2} \frac{\partial^2 \Phi_3}{\partial \xi^2} + \frac{(y+d)^4}{4!} \frac{\partial^4 \Phi_1}{\partial \xi^4}.$$

Surface isobarity:

$$\eta_2 = \frac{c_0}{g} \frac{\partial \Phi_1}{\partial \xi},$$

$$\eta_4 = \frac{c_0}{g} \frac{\partial \Phi_3}{\partial \xi} - \frac{c_0 d^2}{2g} \frac{\partial^3 \Phi_1}{\partial \xi^3} - \frac{1}{g} \frac{\partial \Phi_1}{\partial \tau} - \frac{1}{2g} \left(\frac{\partial \Phi_1}{\partial \xi} \right)^2.$$

Surface impermeability at order ϵ^5 :

$$\frac{\partial \eta_2}{\partial \tau} + \frac{c_0 d^2}{6} \frac{\partial^3 \eta_2}{\partial \xi^3} + \frac{3g}{2c_0} \eta_2 \frac{\partial \eta_2}{\partial \xi} = 0.$$

Korteweg & de Vries equation

Approximation:

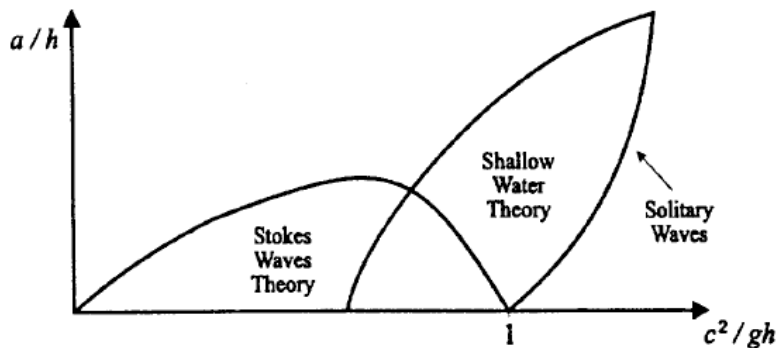
$$\eta \approx \epsilon^2 \eta_2, \quad \frac{\partial}{\partial \xi} = \frac{1}{\epsilon} \frac{\partial}{\partial x},$$
$$\frac{\partial}{\partial \tau} = \frac{1}{\epsilon^3} \frac{\partial}{\partial t} + \frac{c_0}{\epsilon^2} \frac{\partial}{\partial \xi} = \frac{1}{\epsilon^3} \frac{\partial}{\partial t} + \frac{c_0}{\epsilon^3} \frac{\partial}{\partial x},$$

KdV:

$$\frac{\partial \eta}{\partial t} + c_0 \frac{\partial \eta}{\partial x} + \frac{3g}{2c_0} \eta \frac{\partial \eta}{\partial x} + \frac{c_0 d^2}{6} \frac{\partial^3 \eta}{\partial x^3} = 0.$$

Validity of shallow water theories

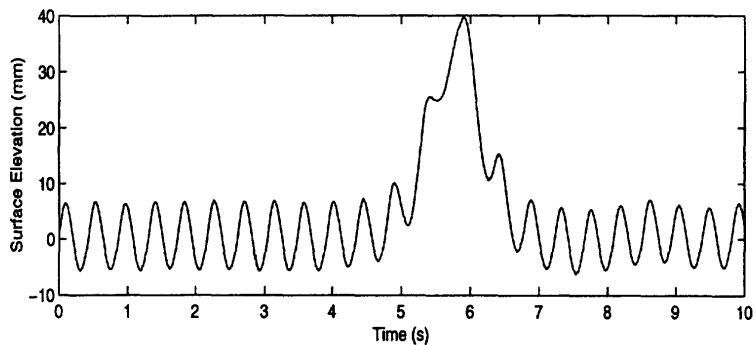
Littman's diagram:



- Stokes expansion: finite radius of convergence.
- Shallow water expansion: divergent series.

Sinusoidal wave over a solitary wave

Experimental measurements:



Hybrid expansion (Hint of the method)

Stokes series:

$$\phi = \sum_{m=1}^{\infty} \sum_{n=-m}^m \epsilon^m A_{m,n}(y) e^{ni(kx - \omega t)}.$$

With 'slow' variables $\xi = \epsilon x$, $\tau = \epsilon t$:

$$\phi = \sum_{m=1}^{\infty} \sum_{n=-m}^m \epsilon^m A_{m,n}(y) e^{ni(k\xi - \omega\tau)/\epsilon}.$$

Nonlinear WKB expansion:

$$\phi = \sum_{m=1}^{\infty} \sum_{n=-m}^m \epsilon^m A_{m,n}(\xi, y, \tau) e^{niS(\xi, \tau)/\epsilon},$$

$$S = \sum_{m=0}^{\infty} \epsilon^m S_m(\xi, \tau).$$

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Amplitude vs Phase

Phase expansion, e.g.:

$$\exp(ni\epsilon S_2) = 1 + ni\epsilon S_2 + O(\epsilon^2).$$

Phase terms can be rewritten as amplitude terms.

Similarly, amplitude terms can be rewritten as phase terms.

Lagrange solution of Laplace equation

Laplace + bottom impermeability:

$$\phi_{xx} + \phi_{yy} = 0 \quad \text{for } y \geq 0, \quad \phi_y = 0 \quad \text{at } y = -d.$$

Lagrange (1791) expansion:

$$\begin{aligned} \phi(x, y, t) &= \sum_{n=0}^{\infty} (y + d)^n f_n(x, t) \quad \Rightarrow \\ \sum_{n=0}^{\infty} (y + d)^n \frac{\partial^2 f_n}{\partial x^2} + \sum_{n=2}^{\infty} n(n-1) (y + d)^{n-2} f_n &= 0 \\ \sum_{n=0}^{\infty} (y + d)^n \left[\frac{\partial^2 f_n}{\partial x^2} + (n+2)(n+1) f_{n+2} \right] &= 0. \end{aligned}$$

with bottom impermeability:

$$\begin{aligned} f_0 &= \phi(x, y = -d, t) \equiv \Phi(x, t), & f_1 &= 0, \\ \phi(x, y, t) &= \sum_{n=0}^{\infty} (-1)^n \frac{(y+d)^{2n}}{(2n)!} \frac{\partial^{2n} \Phi(x, t)}{\partial x^{2n}} \\ &= \cos[(y+d) \partial_x] \Phi(x, t). \end{aligned}$$

If the potential is known at the bottom, it is known everywhere.

If $\Phi \propto \sin(kx - \omega t)$:

$$\begin{aligned} \phi(x, y, t) &= \cos[(y+d) \partial_x] \sin(kx - \omega t) \\ &= \cosh k(y+d) \sin(kx - \omega t) \end{aligned}$$

If $\Phi \propto \tanh(kx - \omega t) \Rightarrow$ complicated (convergent?) series.

General solution of Laplace equation

Laplace equation rewritten:

$$\begin{aligned} \phi_{xx} - i^2 \phi_{yy} &= 0, & i^2 &= -1, \\ \Rightarrow \phi &= F(x + i(y + d)) + G(x - i(y + d)). \end{aligned}$$

Bottom impermeability:

$$\begin{aligned} 0 &= \phi_y(y = -d) = iF'(x) - iG'(x) \\ \Rightarrow G &= F \equiv \frac{1}{2} \phi(y = -d). \end{aligned}$$

Laplace equation + bottom impermeability:

$$\phi(x, y) = \frac{1}{2} \Phi(x + i(y + d)) + \frac{1}{2} \Phi(x - i(y + d)).$$

Application

If $\Phi \propto \tanh(kx)$:

$$\begin{aligned}\phi &= \frac{1}{2} \tanh(kx + ik(y + d)) + \frac{1}{2} \tanh(kx - ik(y + d)) \\ &= \frac{\sinh kx \cosh kx}{\sinh^2 kx + \cos^2 k(y + d)} \\ &= \frac{\tanh kx}{\tanh^2 kx + \operatorname{sech}^2 kx \cos^2 k(y + d)}\end{aligned}$$